

Silicon isotope fractionation during the formation of (amorphous) hydroxyaluminosilicate (HAS) phases

Franziska M. Stamm¹, Andre Baldermann¹, Dorothee Hippler¹ and Martin Dietzel¹

¹Institute of Applied Geosciences, Graz University of Technology, NAWI Graz Geocentre, 8010 Graz

Motivation

- Weathering of silicate minerals is an important process on the Earth's surface
- Key Element: Silicon (Si)
- Critical Zone (Fig. 1):
 - Si is precipitated as short-range order hydroxyaluminosilicate (HAS) solid phase, e.g., allophane (ALO) & hisingerite (HIS)
- Experimental study of Si isotope fractionation between reactive fluids & HAS
- Investigation of kinetic & equilibrium Si isotope fractionation to understand mechanisms of HAS formation

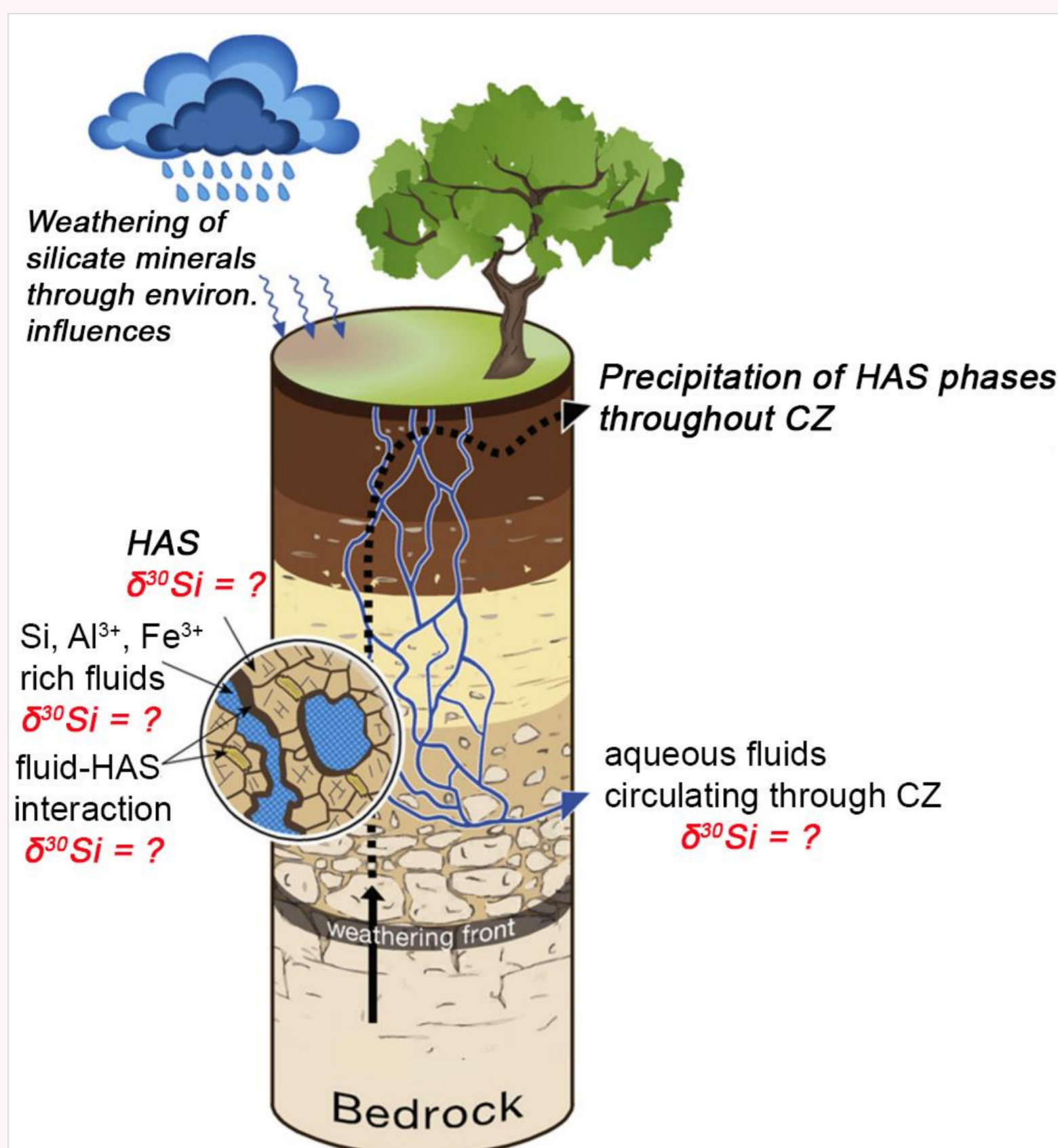


Fig. 1 Model of the Critical Zone (CZ) with processes and pathways leading to HAS formation (modified after v. Blanckenburg et al.)

Nomenclature & Theoretical Background

Si isotopic composition of a phase:

$$\delta^{30}\text{Si}(\text{‰}) = \left(\frac{(^{30}\text{Si}/^{28}\text{Si})_{\text{sample}}}{(^{30}\text{Si}/^{28}\text{Si})_{\text{NBS-28}}} - 1 \right) \times 1000$$

Isotopic fractionation factor between two phases:

$$\Delta_{A-B} = \delta_A - \delta_B \approx 10^3 \ln \alpha_{A-B}$$

Three isotope method (Fig. 2)^{2,3}

- Determination of equilibrium isotope fractionation factors
- Precise determination of exchange reactions of small extent
- Monitoring of reaction kinetics
- Isotopic exchange between 2 phases:
 - Phase A natural composition: plots on terrestrial fractionation line (TFL)
 - Phase B: enriched in ²⁹Si
- Isotope exchange complete (100%) on secondary fractionation line (SFL)

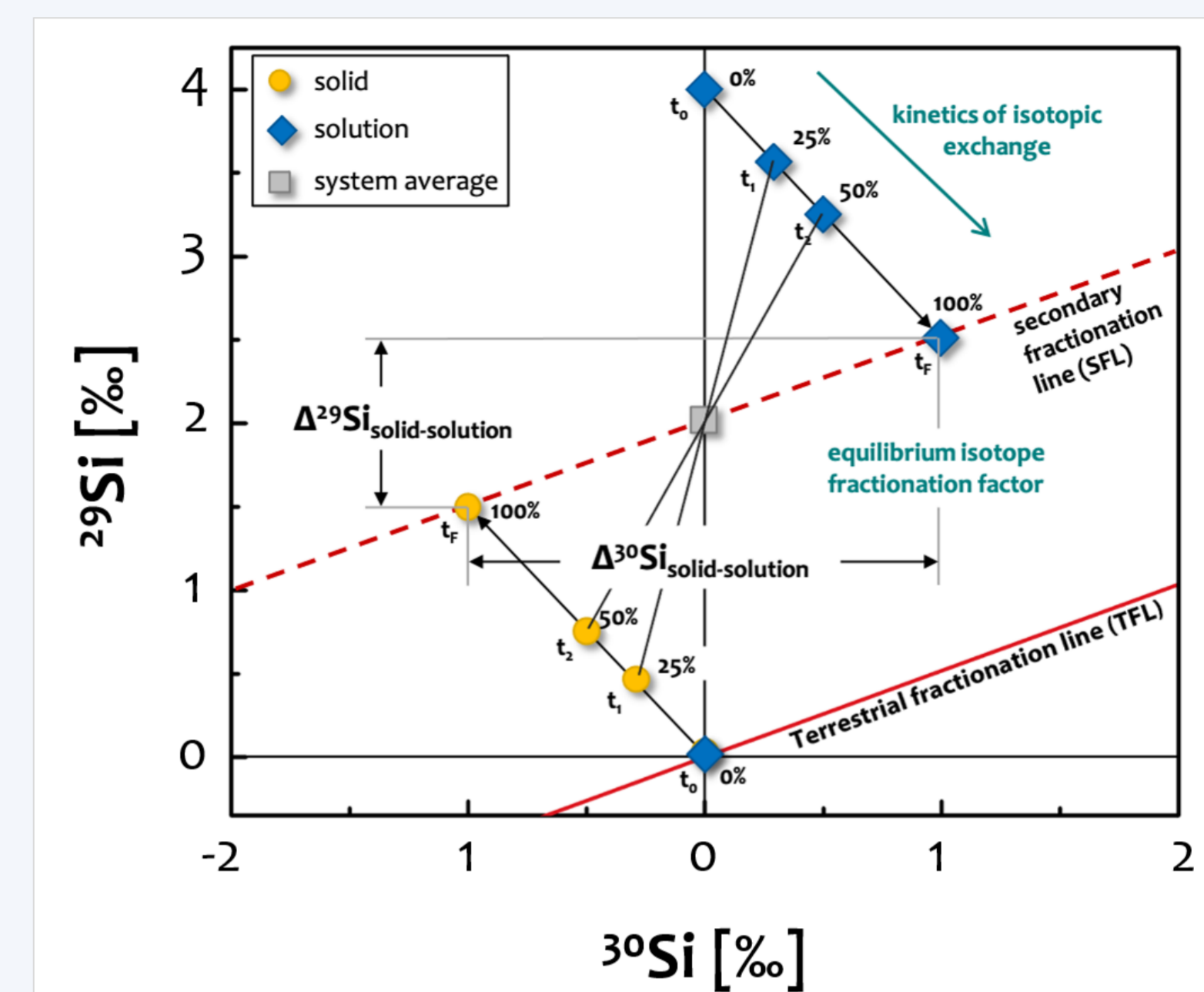


Fig. 2 Schematic diagram of the three-isotope method applied to Si. modified after Stamm et al.³

Experimental Setup

Starting Solutions

- Si(OH)₄ rich solution prepared from Na₂SiO₃ · 5H₂O (pH ~13)
- Al³⁺ and Fe³⁺ solutions prepared from AlCl₃ · 6H₂O and FeCl₂ · 5H₂O (pH~2.7)

Experiments @ 25°C

Step 1: Precipitation experiments

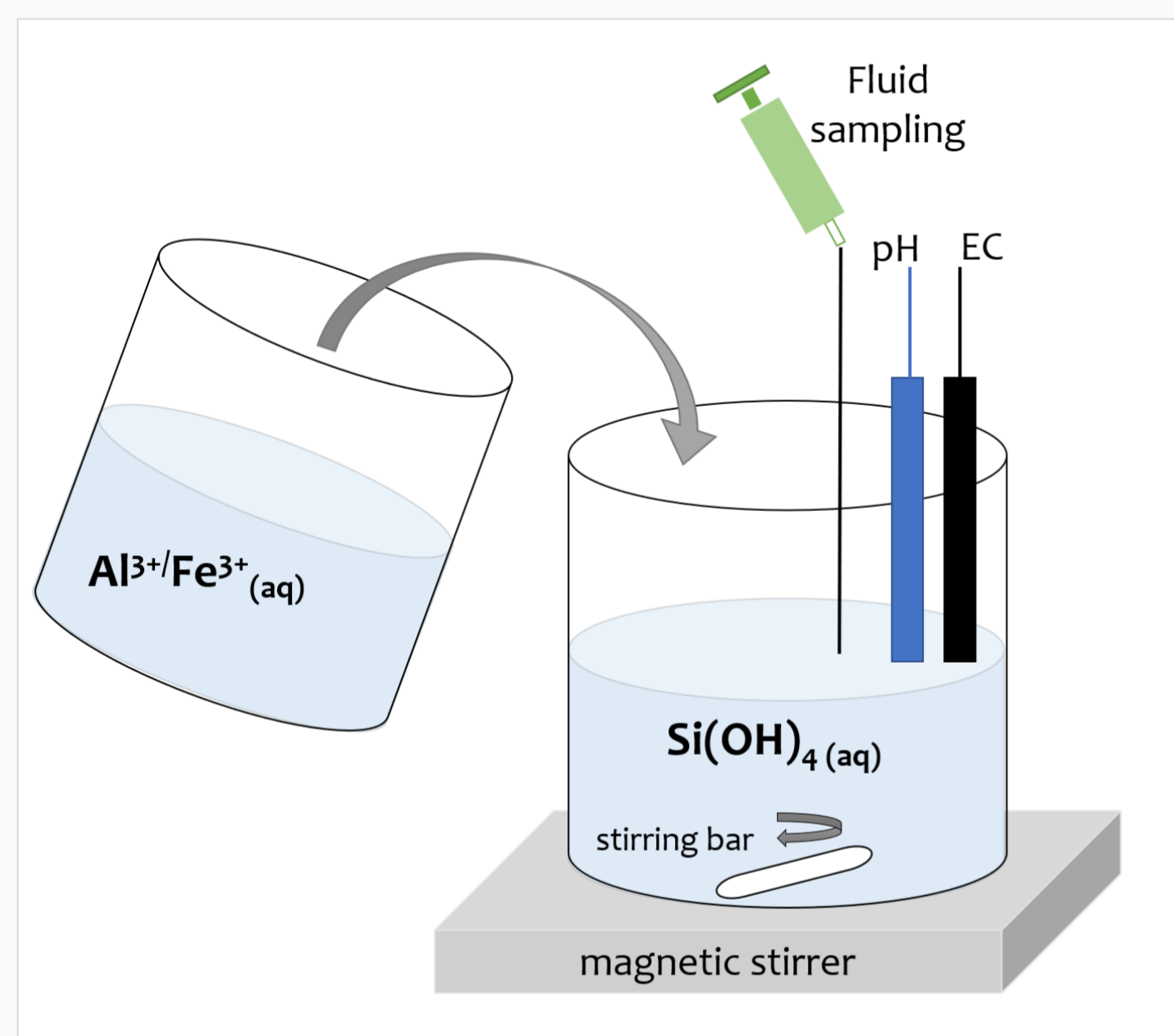


Fig. 3 Schematic diagram of the precipitation experiments of two HAS phases: allophane and hisingerite (Al or Fe:Si=1:1; Ratz et al.⁴)

- Time series sampling to monitor kinetic isotope effects during precipitation of allophane (~Al₂³⁺O₃(SiO₂)_{1.3-2} · H₂O_{2.5-3}; ALO) / hisingerite (~Fe₂³⁺Si₂O₅(OH)₄ · 2H₂O; HIS)

Step 2: Isotope exchange experiments

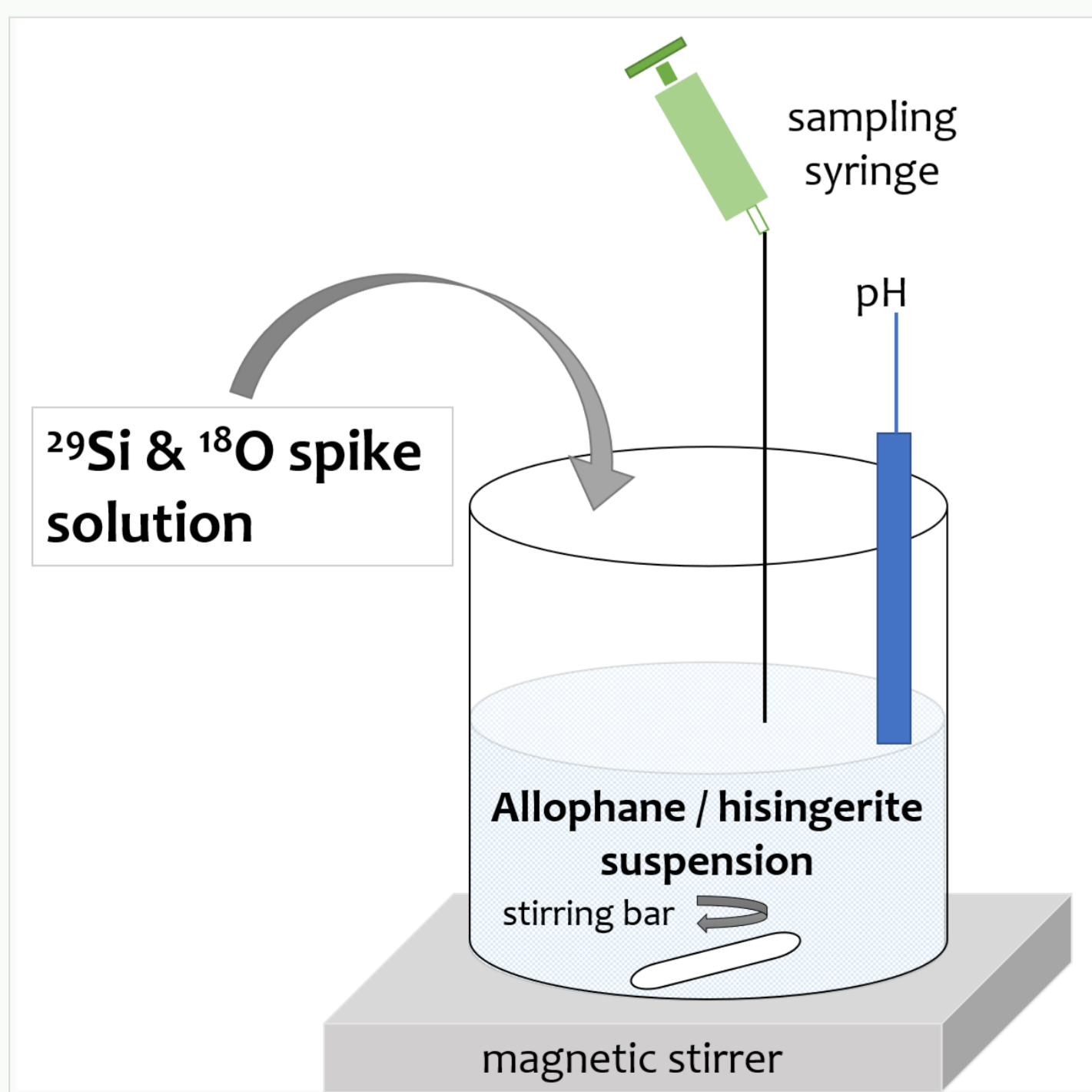


Fig. 4 Schematic diagram of isotope exchange experiments (modified from Ratz et al.⁴)

- Attainment of **chemical equilibrium**
- Spiking of equilibrated solution with ²⁹Si & ¹⁸O
- Isotope exchange experiments
 - Closed system
 - Time-series sampling

Preliminary Results

Isotopic Exchange

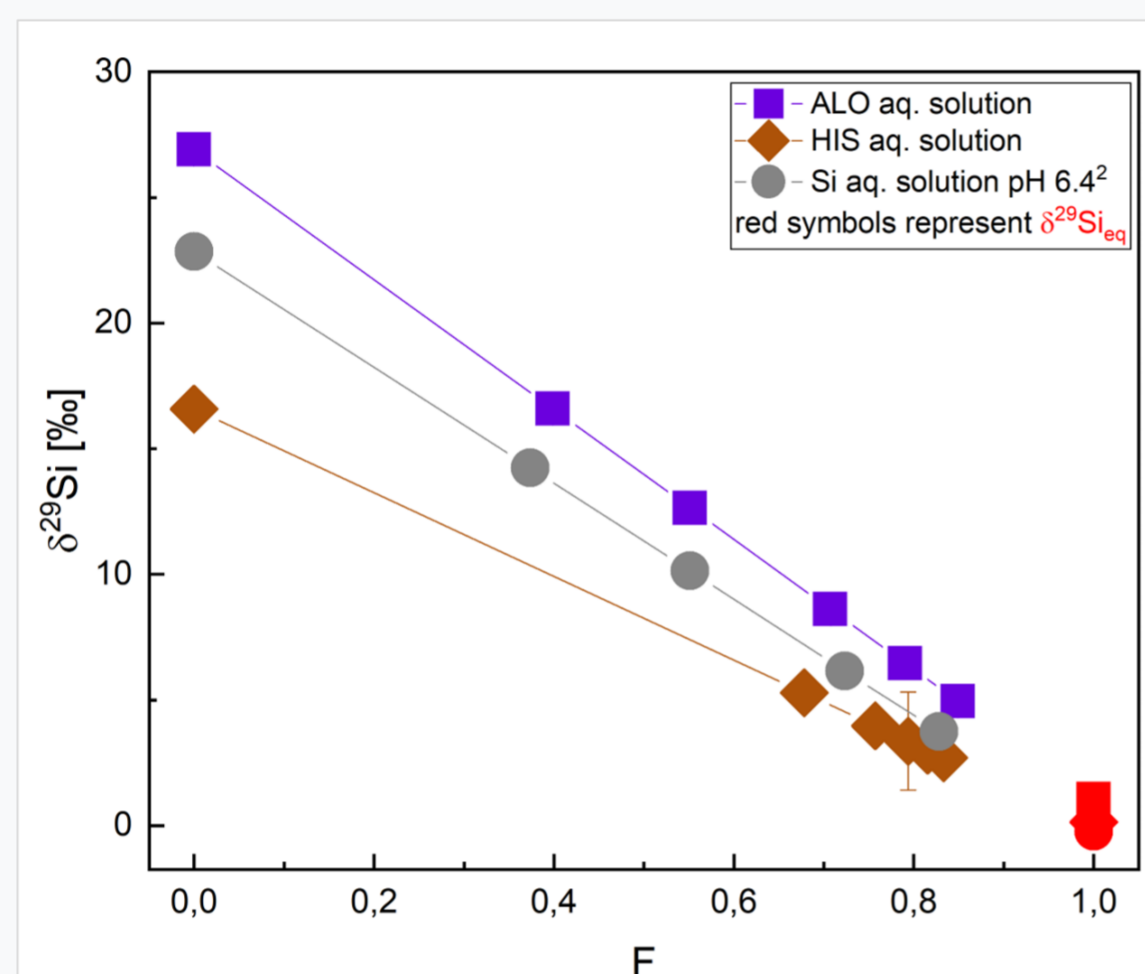


Fig. 5 Evolution of δ²⁹Si over the degree of isotopic exchange (F); 2SD shown within the symbols.

- δ²⁹Si_{aq}-F diagram shows fast depletion in δ²⁹Si in solution within 15 days
- Allophane (ALO) & hisingerite (HIS) show similar behaviour

- Degree of isotopic exchange (F) determined for any reaction time t:

$$F = (\delta_t - \delta_i) / (\delta_e - \delta_i)$$

δ_t - δ at reaction time t of isotopic exchange
δ_i - initial δ of solution
δ_e - δ at equilibrium

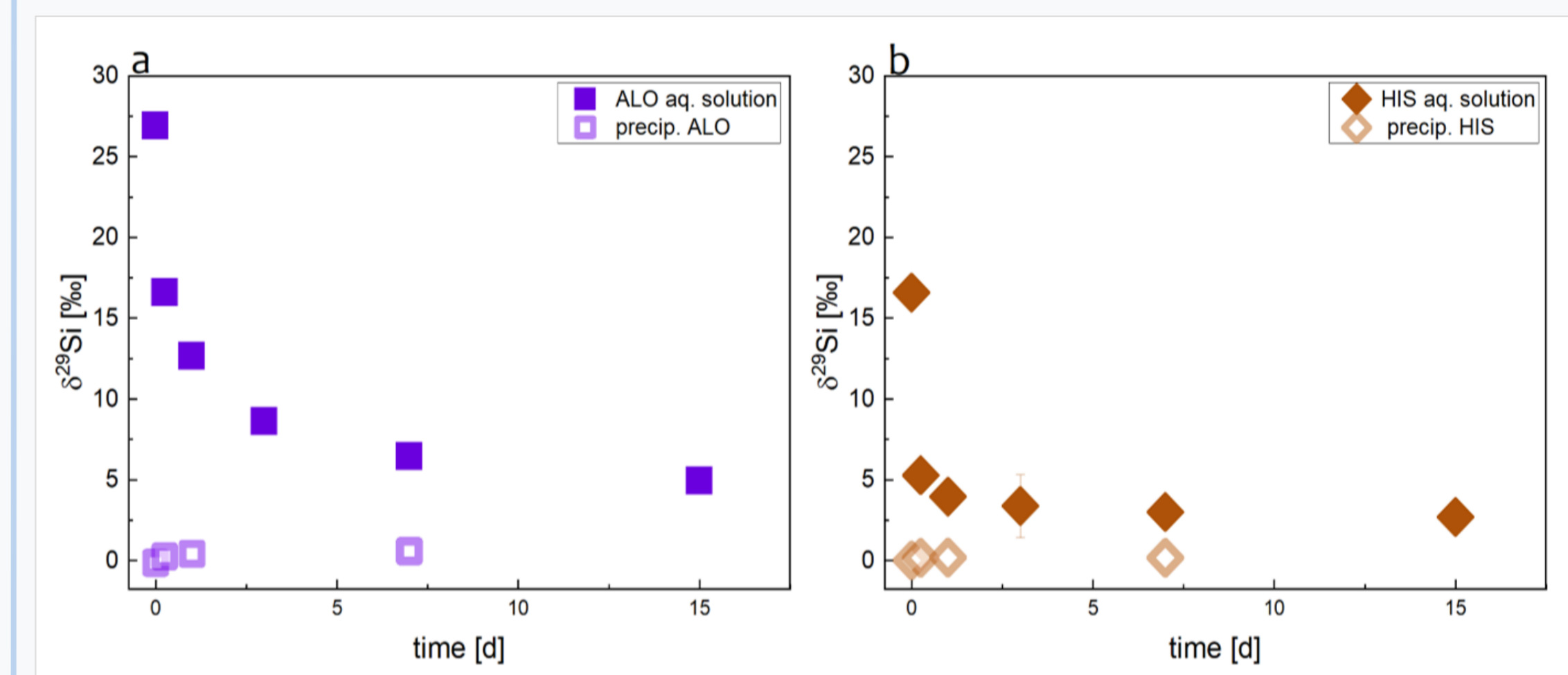


Fig. 6 Evolution of δ²⁹Si over time for the isotope exchange experiments with a) allophane (ALO) and b) hisingerite (HIS); 2SD if not shown, within the symbols

- Evolution of δ²⁹Si of the solid and reactive solution over reaction time throughout allophane (ALO) and hisingerite (HIS) formation
- Δ³⁰Si_{eq} over reaction time indicates clear and fast depletion of spiked element

Determination of Δ³⁰Si_{solid-Si(aq)}

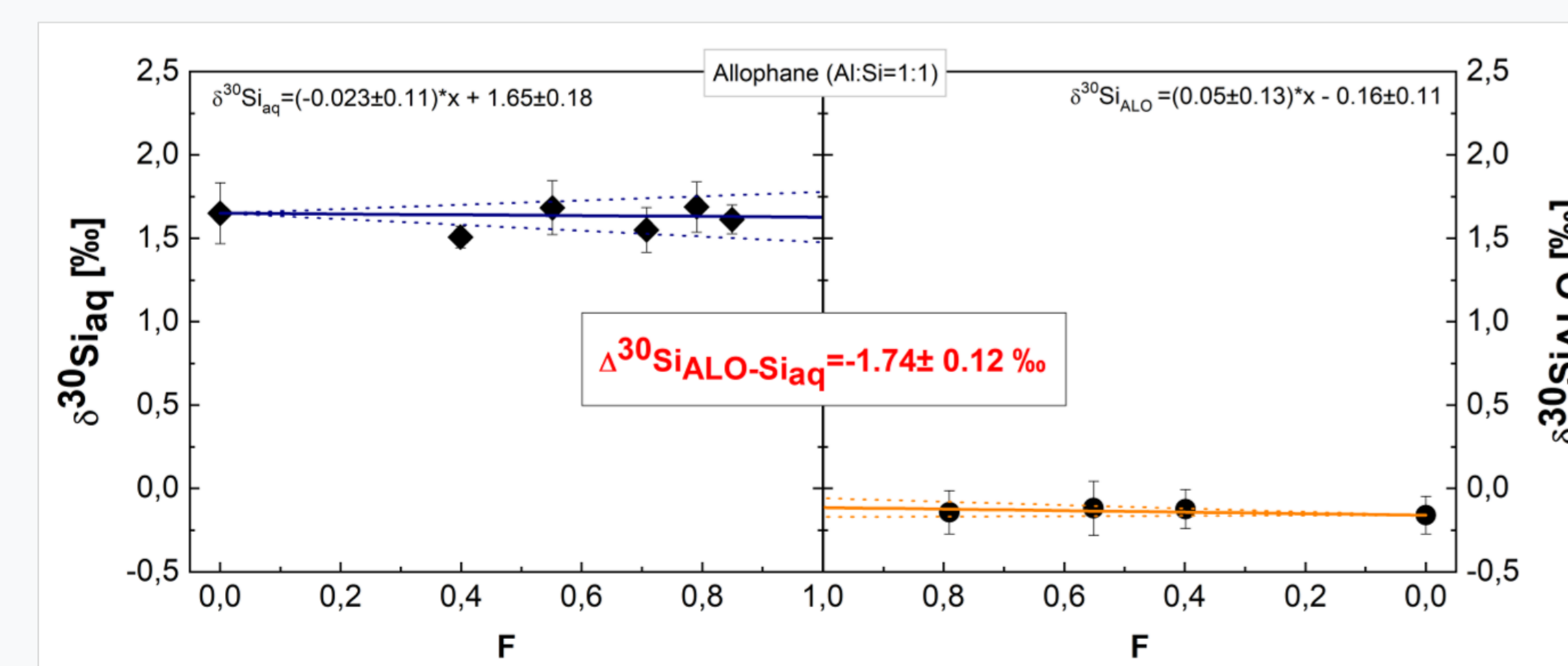


Fig. 7 Determination of the equilibrium fractionation factor (Δ³⁰Si_{ALO-Si(aq)}) between allophane and the reactive solution by extrapolation (dashed line 95% confidence interval).

- Δ³⁰Si_{eq} between solids (allophane/ hisingerite) and aqueous solutions determined by extrapolation of δ³⁰Si isotopic compositions of the two phases

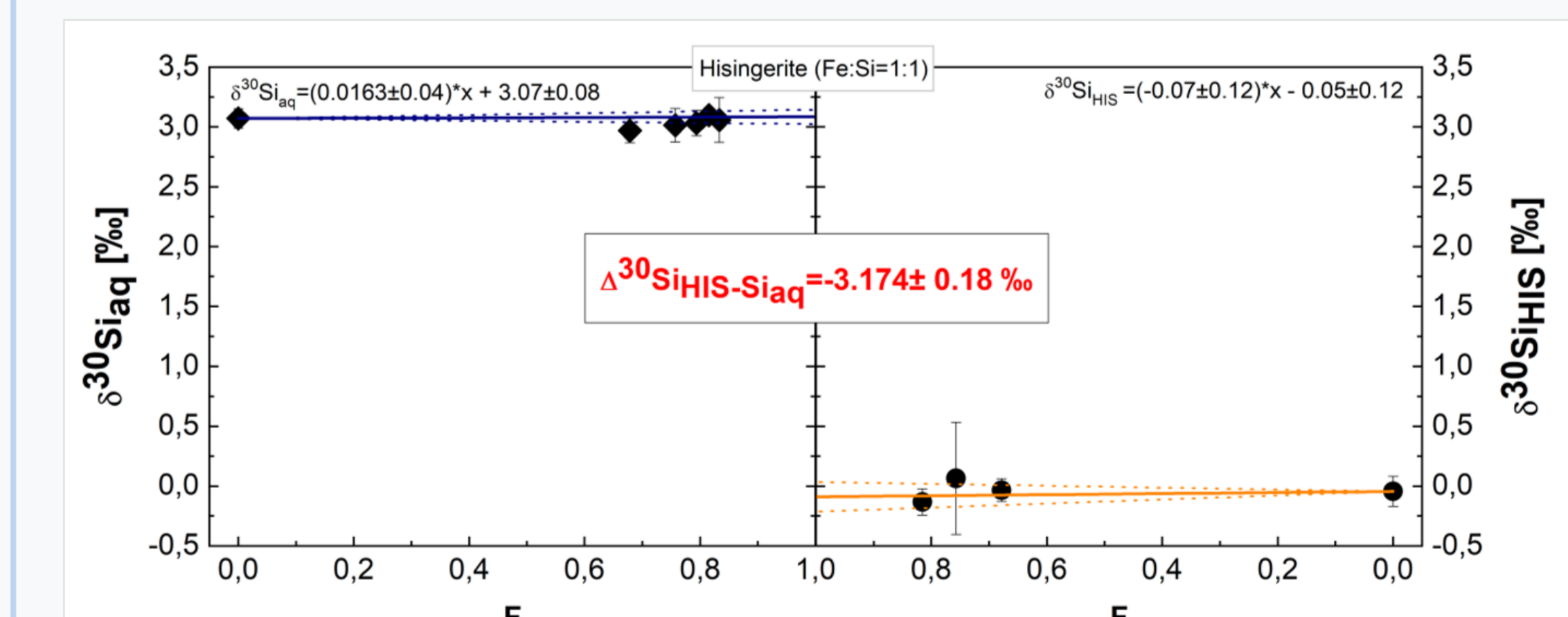


Fig. 8 Determination of the equilibrium fractionation factor (Δ³⁰Si_{HIS-Si(aq)}) between hisingerite and the reactive solution by extrapolation (dashed line 95% confidence interval).

Conclusions & Outlook

- Light isotopes are preferentially incorporated in the HAS precipitates
- Si isotope re-distribution between allophane/ hisingerite and reactive solutions occurs within days
 - highly dynamic exchange between solid and aqueous Si tetrahedrons
- Si isotope exchange kinetics slightly increase in the order of HIS>ALO & pure Si system
 - Polymerization steps more relevant than Me-O-Si for isotopic re-equilibration?
- Outlook**
 - Study of the evolution of the isotopic composition during allophane & hisingerite precipitation
 - O isotope exchange experiments to investigate coupling between Si and O isotope exchange reaction mechanisms
 - Determination of Si and O isotope exchange kinetics and apparent/equilibrium fractionation factors
- Implications: assessment and monitoring of inverse weathering scenarios by (coupled) Si and O isotope distribution between interstitial solutions and HAS precipitates in natural surroundings

References & Acknowledgements

[1] v. Blanckenburg et al. (2012) EPSL, 351-352, 295-305; [2] Matsuhashi et al. (1978) GCA 42, 173-182; [3] Stamm et al. (2019) GCA 255, 49-68; [4] Ratz et al. in prep.
Special thanks go to Bettina Ratz (TU Graz), Daniel A. Frick, Josefine Holtz, and Friedhelm v. Blanckenburg at the GFZ – German Research Centre for Geosciences (Helmholtz Centre Potsdam)